GROUNDWATER IN THE BUSSELTON AREA, PROGRESS REPORT ON EXPLORATORY DRILLING

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INTRODUCTION

Between September 1966 and August 1967 an east-west line of exploratory groundwater bores was drilled in the Busselton area in the southern part of the Perth Basin. The bores were drilled to investigate groundwater occurrence in the rapidly developing coastal areas between Busselton and Dunsborough and also as part of the general investigation of the Perth Basin. Five bores ranging from 1,485 feet to 2,010 feet deep were drilled at locations shown on Plate 1.

The drilling confirmed the presence west of Busselton of major faulting beneath a strong unconformity seperating underlying older sediments of Permian and Lower Jurassic ages from an overlying blanket of South Perth Formation and Quaternary sediments. The faulting divides the western section of the basin into fault blocks of different ages. Groundwaters in the sediments west of Quindalup bore 1 are generally of poor quality and supply, and unsuitable for domestic use or irrigation. In Quindalup bores 1, 4, and 5 promising supplies of low salinity groundwater were encountered in Lower Jurassic to Lower Cretaceous sediments. This water is suitable for domestic use, but because of high bicarbonate content, some may have to be treated before being acceptable for irrigation use on clayey soils.

The bores were drilled using the Mines Department Mayhew 2000 rotary rig, with sludge samples collected every 10 feet and cores taken wherever practicable. Water samples were collected by means of Halliburton and Johnson Formation Testers and samples sent for analysis at the Government Chemical Laboratories. The bores were geophysically logged by the Geological Survey and routine palynology was carried out on selected samples. The results of chemical analyses and palynological determinations are shown in Appendices 2 and 3.

PHYSIOGRAPHY

The southern part of the Perth Basin has been divided (Lowry, 1965) into four broad physiographic regions.

The Swan coastal plain is a flat dune, soil, and alluvium covered plain extending northwards along the coast of Geographe Bay, bounded in the south by the Whicher scarp, and in the east by the Darling fault scarp. The plain has an average width of about 12 miles.

The Blackwood area is in the central section of the basin, and has a maximum elevation of about 500 feet, capped by sand and laterite and underlain by Mesozoic sediments.

The Leeuwin—Naturaliste ridge, which marks the western margin of the basin, has a maximum elevation of about 700 feet and is a north-south trending ridge of Precambrian crystalline rocks capped by laterite and Coastal Limestone. The ridge is bounded on the east by the Dunsborough fault.

The Scott coastal plain is a low lying swampy flat plain marked by scattered dune ridges.

GENERAL GEOLOGY

The southern part of the Perth Basin consists of a deep trough of sediments flanked on the east and west by Precambrian crystalline rocks of the Western Australian Shield and the Leeuwin-Naturaliste ridge. The trough continues to the north and south beneath the Indian Ocean, forming a graben between the Dunsborough and Darling faults.

Geophysical evidence suggests that a maximum thickness of 15,000 to 20,000 feet of easterly dipping sediments is present in the central and eastern sections of the basin, and that this is

intersected by four major, branching, north-south trending, faults having throws of up to 8,000 feet.

Lowry has described in detail the surface geology, which consists of a few scattered outcrops of Mesozoic sediments mainly masked by Quaternary deposits including alluvium, dune sands, and Coastal Limestone in the coastal plains, and sand and laterite in the Blackwood area.

Bores have intersected the sediments underlying the Quaternary sequence but mostly they are shallow and have been confined to the Lower Cretaceous—Upper Jurassic South Perth Formation, which probably blankets older sediments over a large proportion of the basin. However Sue No. 1, Alexandra Bridge No. 1 petroleum exploration wells and several coal bores in the southwest encountered Lower Permian sediments beneath the South Perth Formation, and the Abba River and Western Titanium bores encountered the Upper Jurassic Yarragadee Formation.

Quindalup bores 1 to 5

Drilling in the present part of the programme was confined to the Swan coastal plain. Bores 1 to 3 were located on the fringes of Geographe Bay 2½, 6½ and 10 miles west of Busselton respectively, and bores 4 and 5 at Wonnerup and Ruabon, 8 and 12 miles east of Busselton (Plate 1).

The bores intersected a thin veneer of Quaternary sediments overlying South Perth Formation, which in turn is separated from underlying older sediments by a strong unconformity. These sediments include those of Permian, Lower Jurassic, and Upper Jurassic ages.

The stratigraphic relationship between the bores is shown in Table 1.

Quaternary

The Quaternary sediments vary from 20 to 35 feet in thickness, forming a veneer above the South Perth Formation. In bores 1 to 3 they consist of well-sorted, fine-grained sand, marly clay, and shelly limestone of the Quindalup Dune System. In bores 4 and 5 farther inland, the sediments are poorly sorted, iron-stained sand and shelly limestone which possibly belong to the older Bassendean Dune System.

Lower Cretaceous to Upper Jurassic

Current thought on the South Perth Formation includes sediments described by previous workers as belonging to the Capel River Group and Yarragadee Formation.

In the Quindalup bores the thickness ranges from 270 feet in bore 4 to 1,180 feet in bore 5. The sediments are dominantly of continental type but rare microplankton indicate some marine influence during deposition. They are typically poorly consolidated, finely current-bedded, and lenticular. Correlation between bores is difficult and suspect, but the general dip is very gentle from east to west.

The sequence consists mainly of silt, siltstone, clay, poorly-sorted sand and sandstone containing abundant carbonaceous material, pyrite, and mica, and thin bands of lignite.

The arenaceous sediments are feldspathic, and are more consolidated with depth, making up 60 to 75 per cent of the total section.

The South Perth Formation unconformably overlies older sediments in each of the Quindalup bares

Upper Jurassic

The Upper Jurassic Yarragadee Formation was encountered in Quindalup bores 4 and 5 and in nearby bores in the Abba River and Capel areas where it underlies the South Perth Formation.

he section is dominantly an arenaceous contial deltaic deposit containing less than 20 per clay and silt. The sandstone and sand differ from those of the overlying South Perth Foron but are typically garnetiferous and slightly consolidated. They are generally moderately porly sorted, medium to very coarse in grain, slightly feldspathic. Interbedded siltstone, clay, stone, and rare lignite make up the remainder.

r Jurassic

wer Jurassic sediments of continental type are not between 630 and 1,930 feet in Quindalup 1. They consist of cross-bedded, poorly-sorted, to coarse-grained sand and poorly consolidated stone interbedded with mottled grey to red tone, grey to green pyritic shale and siltstone. claystone shows slickensiding and slumping.

erlain unconformably by the South Perth Foron, the Lower Jurassic material in Quindalup 1 is the first observed occurrence of sediments is age in the southern part of the Perth Basin, s equated to the Cockleshell Gully Formation.

0.00

mian sediments, also overlain unconformably e South Perth Formation, occurred in Quindaores 2 and 3, the westernmost bores of the

Quindalup bore 2 they are present between nd 1,807 feet, and consist of well consolidated to coarse-grained, slightly argillaceous sand-containing abundant garnet, finely currented, grey siltstone, 17 seams of carbonaceous, and poor quality dull coal ranging from 7 to ct in thickness. The sandstone contains rare reous material.

Quindalup bore 3 the sequence is lithologically ar except that only one 8-feet thick dull coal was encountered, although carbonaceous rial commonly occurred as partings on bedplanes.

e Permian sequences are of different ages, that lindalup bore 3 being the younger of the two.

STRATIGRAPHY AND STRUCTURE

stratigraphic relationships are shown in 2, and in Table 1.

Quaternary material forms a thin veneer the South Perth Formation, and the maximum and thickness of 35 feet is probably exceeded in the Wonnerup area where gamma-ray logof existing bores indicates a thickness of 50 to 120 feet near the Wonnerup Estuary.

e South Perth Formation is present throughthe drilling area and ranges widely in thickness
345 to 668 feet west of Busselton, thickening
er 1,000 feet in the Milne Street, Abba River
of Quindalup 5 bores. It thins again to 180
of feet in the Wonnerup-Capel area. The
section between Abba River 1 bore and Capel
of 2) shows this variation, which may either
to deposition over a deeply dissected Upper
sic surface or to post-Yarragadee Formation
of There is little concrete evidence to suphe latter supposition, but differences in levels

of basalt flows encountered in Quindalup bore 5 and Abba River bore 3 may either be due to depositional dip or to minor faulting. If a deep trough of South Perth sediments exists it probably is an old erosional channel which may extend as far east as the Whicher scarp, and possibly be related to basalt occurrences in the area.

The variation in thickness of the South Perth Formation illustrates the strong unconformity which separates these beds from the underlying strata. Below the unconformity three major northsouth trending faults with throws of several thousand feet have been proved by the drilling. These include the Wirring and Busselton faults inferred from previous gravity and seismic surveys, and a fault postulated by Lodwick from seismic data. The faults divide the area beneath the unconformity west of Busselton into four blocks containing Upper to Lower Permian, Lower Permian, Lower Jurassic, and Upper Jurassic sediments. The block between the Wirring and Busselton faults is uplifted with respect to the others, which are normal east block down.

The Yarragadee Formation is present east of Busselton in Quindalup bores 4 and 5, Abba River bores 1 and 3, and in bores at the Western Titanium mining pit at Capel. No bores have penetrated the full succession of the Jurassic sediments which reach a maximum observed thickness of 1,615 feet in Quindalup bore 4.

Lower Jurassic sediments in Quindalup bore 1 are of similar lithology and age to the Cockleshell Gully Formation in the Hill River area, but the arenaceous sediments appear to be generally coarser in grain. The Lower Jurassic sediments are restricted in occurrence to the fault block immediately west of Busselton and nothing is yet known of their extension to the south.

Lower Permian coal-bearing sediments in Quindalup bore 2 are probably Upper to Lower Artinskian in age. The lower part of the section contains similar spore assemblages to the lower Artinskian coals of the Collie and Irwin River areas and to those encountered at 8,866 feet in WAPET's Sue No. 1 well. In this part of the basin these sediments are confined to the wedge-shaped block between the Wirring and Busselton faults.

The Permian sediments in Quindalup bore 3 appear younger than those in bore 2 and have been tentatively dated as early Upper Permian to Upper Artinskian. They lie between the Dunsborough and Wirring faults and their southward extent is not known.

OCCURRENCE OF GROUNDWATER

Groundwater was present in sediments of all ages and was sampled wherever possible by formation testers. The water samples were used by D. L. Rowston as controls for the estimation of salinities from resistivity logging. These results are shown on bore completion reports and on Plate 1.

Quaternary aquifers

Small supplies of groundwater are available from the thin veneer of Quaternary sediments. The water varies from stock to domestic quality and

TABLE 1
STRATIGRAPHIC SUCCESSION IN QUINDALUP BORES (EAST TO WEST)

| Age | Formation | Bore 3 (feet) | Bore 2 (feet) | Bore 1 (feet) | Bore 4 (feet) | Bore 5 (feet) |
|---------------------|---|-------------------|-------------------------|---------------|---------------|---------------|
| 4 | Recent sediments and Quindalup and Bassendean Dune Systems | 0-32 | 0-20 | 0-22 | 0-35 | 0-22 |
| Cretaceous to Upper | South Perth Formation | 32-700 | 20-365 | 22-630 | 35-305 | 22-1180 |
| uzassic | Yarragadee Formation | Not present | Not present | Not present | 305–1920 | 1180-2010 |
| waste | Cockleshell Gully Sandstone equiv- alent | Not present | Not present | 630-1930 | | 17.11 |
| pjet Permian | | 700 to 1485 | Not present 865-1807 | | | |

Correlation based on lithology, palynology, and geophysical data.

usually contains high $\rm H_2S$ and bicarbonate. Bores near the coast and tidal drainage channels are subject to saline intrusion. The sediments form a useful source of domestic and garden water in coastal cottages in and around Busselton.

South Perth Formation aquifers

The aquifers of the South Perth Formation are lenticular and contain groundwater of variable quality and supply. Lowry (1965) has summarised bore census data for the Busselton area and gives the following relationship between bore depth and frequency:

Table 2

| | Bore | depth | (feet) | (average | 60) | No. of bores |
|---|------|-------|--------|----------|-----|----------------------------------|
| 31-50 51-70 71-100 101-150 >150 | | | | | | 123 123 66 37 10 |

These sediments are the most important source of groundwater for farmers, particularly the shallow glauconitic sands near Quindalup. Geneshallow glauconitic sands near Quindalup. Generally bores are sited for convenience, and as the beds are very lenticular the landholder continues drilling until he obtains water. Most supplies are sub-artesian, yielding 200 to 1,000 gallons per hour, and range from good stock to domestic quality, although saline groundwater has been encountered in some areas west of Busselton.

West of Busselton on the fringes of Geographe Bay, groundwater in the South Perth Formation of Quindalup bores 1 to 3 is generally of poor quality. In bore 1 thin layers of stock quality groundwater are intercalated with layers ranging in salinity from 8,000 to 20,000 ppm TDS. In bore 2 the quality is uniformly poor, avering about 14,000 ppm TDS. Farther west the quality improves to 450 to 1,500 ppm TDS.

Between Quindalup bore 1 and Capel, domestic Between Quindalup bore 1 and Capel, domestic quality groundwater is obtained at depths of 50 to 1,000 feet from the South Perth Formation. A little saline groundwater has been met near the coast at shallow depths but generally the supply and quality improves with depth. Supplies are generally good ranging from 200 gph in shallow bores to over 30,000 gph in bores 500 to 1,000 feet deep. Near the coast the bores flow.

Yarragadee Formation aquifer

Yarragadee Formation aquifer
Sediments of the Yarragadee Formation were encountered 305 and 1,920 feet in Quindalup bore 4, and between 1,180 and 2,010 feet in Quindalup bore 5. They have also been encountered in the Abba River bores 1 and 3 and in the Western Titanium works at Capel. In the Wonnerup area the sequence is dominantly arenaceous and formation testing suggests that large quantities of domestic quality groundwater (200 to 400 ppm TDS) are available. In the Capel area, bores produce 20,000 to 25,000 gallons per hour from similar but more argillaceous beds. In the Abba River area, poor results obtained in drilling the three Abba River bores probably reflect bad drilling techniques. ling techniques.

Cockleshell Gully Formation (?) aquifers

The Lower Jurassic sand and sandstone in Quindalup bore 1 contain domestic quality ground-water between 670 and 720 feet and between 952 and 1,520 feet. Between 720 and 950 feet and below 1,520 feet the salinity increases to about 2,000 ppm TDS 2,000 ppm TDS.

Formation testing of these sands indicates that they contain good supplies which are possibly artesian.

Permian aquifers

The well consolidated Upper and Lower Permian sandstones in Quindalup bores 2 and 3 vary greatly in porosity and permeability. They contain groundwater ranging from good to poor stock quality (1,300 to 8,800 ppm TDS) and are probably of poor supply.

CHEMICAL CHARACTERISTICS OF THE GROUNDWATER

Standard chemical analyses were obtained for most of the deeper bore samples in the Busselton-Capel area, and the results are shown in Appendix 3. Appendix 4 shows calculated sodium adsorption calcium—magnesium ratios, and residual sodium bicarbonate.

In the Busselton—Capel area most groundwater has a salinity of less than 600 ppm TDS and only west of Busselton, in the South Perth, Lower Jurassic, and Permian formations, do some salinities exceed this figure.

The two main groundwater sources, the South The two main groundwater sources, the South Perth and Yarragadee Formations, contain low salinity water of very similar characteristics. They have a pH ranging from 6.3 to 7.6, contain high concentrations of dissolved iron, and are typically low in calcium, magnesium, and fluoride ions. Variation lies mainly in the bicarbonate—chloride ion content of the South Perth Formation aquifers. These show a decrease in the bicarbonate—chloride. These show a decrease in the bicarbonate—chloride ratio from east to west, which appears to be related to the incidence of old calcareous dune systems on the Swan coastal plain. The calcium—magnesium ratio shows little variation although there is a general trend of increase from east to west, indicating westerly groundwater movement.

The Yarragadee Formation aquifers show primary salinity, and both calcium-magnesium and bicarbonate-chloride ratios increase from east to west across the basin, indicating groundwater movement to the west. The bicarbonate ion content place decreases with death and is possibly related movement to the west. The bicarbonate ion content also decreases with depth and is possibly related to intake from overlying dune areas. The bicarbonate content is, with one possible exception, marginal to high, and untreated groundwater is generally unsuited for use in the irrigation of clayey soils. The exceptional groundwater, which probably belongs to the Yarragadee aquifer system, was in Yoganup bore 1 in the eastern section of the Swan coastal plain at the base of the Whicher scarp. This water has a low pH and very low bicarbonate content and suggests that groundwaters at the eastern margin of the coastal plain near the Whicher scarp may be more suitable for irrigation use. irrigation use.

RECHARGE

Only limited data are yet available on static water levels. Between Busselton and Yoganup two distinct water levels are present in the Yoganup distinct water levels are present in the logarity bores and Quindalup bores 4 and 5, suggesting different potentiometric surfaces for the South Perth and Yarragadee aquifer systems. These have average downward gradients from east to west of 4 and 2½ feet per mile respectively over most of the value but his towards the Whicher scarn plain, but rise towards the Whicher scarp.

Data from levels, lithology, and chemical analyses suggest that recharge of the South Perth aquifers is mainly by direct infiltration of rainfall and river water from the Swan coastal plain. The recharge is selectively governed by topography and the lenticular nature of the sediments, particularly in the areas west of Busselton where saline groundwater is present in both the South Perth Formation and Permian strata. Formation and Permian strata.

The Yarragadee Formation probably receives groundwater from the elevated areas between the Darling fault and the Swan coastal plain, although some direct recharge may occur if areas of shallow Upper Jurassic sediments are present in the eastern sections of the plain.

AVAILABILITY AND USE OF GROUNDWATER

At present a maximum of about 10 million gallons per day of groundwater is extracted from the rons per day of groundwater is extracted from the Yarragadee and South Perth aquifer systems for stock, domestic, town supply, and industrial purposes in the Busselton-Capel areas. It is not possible to estimate the total amount of groundwater available in this area but probably 20 million gallons could be extracted safely without depleting supplies. supplies.

All low salinity waters are suitable for domestic use, but the waters of the Jurassic and deeper south Perth aquifers contain moderate concentrations of bicarbonate ions which would make them unsuitable for the irrigation of heavy clay soils and use in some industries without treatment. Appendix 4 shows the relationship of sodium adsorption ratios and residual bicarbonate (Eaton, 1950). Calculations based on Handa (1964) indicate that the groundwater could be made suitable for irrigation by the application of between 0.2 and 0.6 tens of gypsum per acre foot of water used.

CONCLUSIONS AND RECOMMENDATIONS

The drilling has shown that large supplies of low salinity groundwater are available from Lower Cretaceous and Upper Jurassic sediments east and north of Busselton. This groundwater is suitable for domestic and some industrial use, but as usual in groundwaters from this area, has a high iron content. Some of the groundwaters also contain moderate concentrations of bicarbonate ions which make them unsuitable for use in irrigation without treatment. This should be no great problem.

treatment. This should be no great problem.

West of Busselton the groundwater in sediments underlying the near coastal strip is of marginal to very poor quality, and between Busselton and the Busselton fault, moderate to low salinity groundwater suitable for domestic use was encountered only in the Lower Jurassic sediments. Small quantities of domestic to stock quality groundwater are present in shallow Quaternary sands or in low lying areas underiain by Lower Cretaceous sediments east of Geographe Bay. These supplies are too small and irregular to be of great value other than for farming purposes.

Plate 1 shows the location of five drill-holes proposed for the second stage of the programme. Four bores continue in a line between Ruabon and Donnybrook and will investigate the stratigraphy and hydrological potential of the eastern section of the Swan coastal plain and the elevated area of Mesozoic sediments between Yoganup and Donnybrook. A further bore is proposed to investigate the presence of groundwater between the Sabina and Vasse Rivers, 5 miles southeast of Busselton, where poor and conflicting results were obtained during the drilling of the Abba River bores in 1958.

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Appendix 1

LIST OF DEEP BORES IN THE BUSSELTON—CAPEL AREA

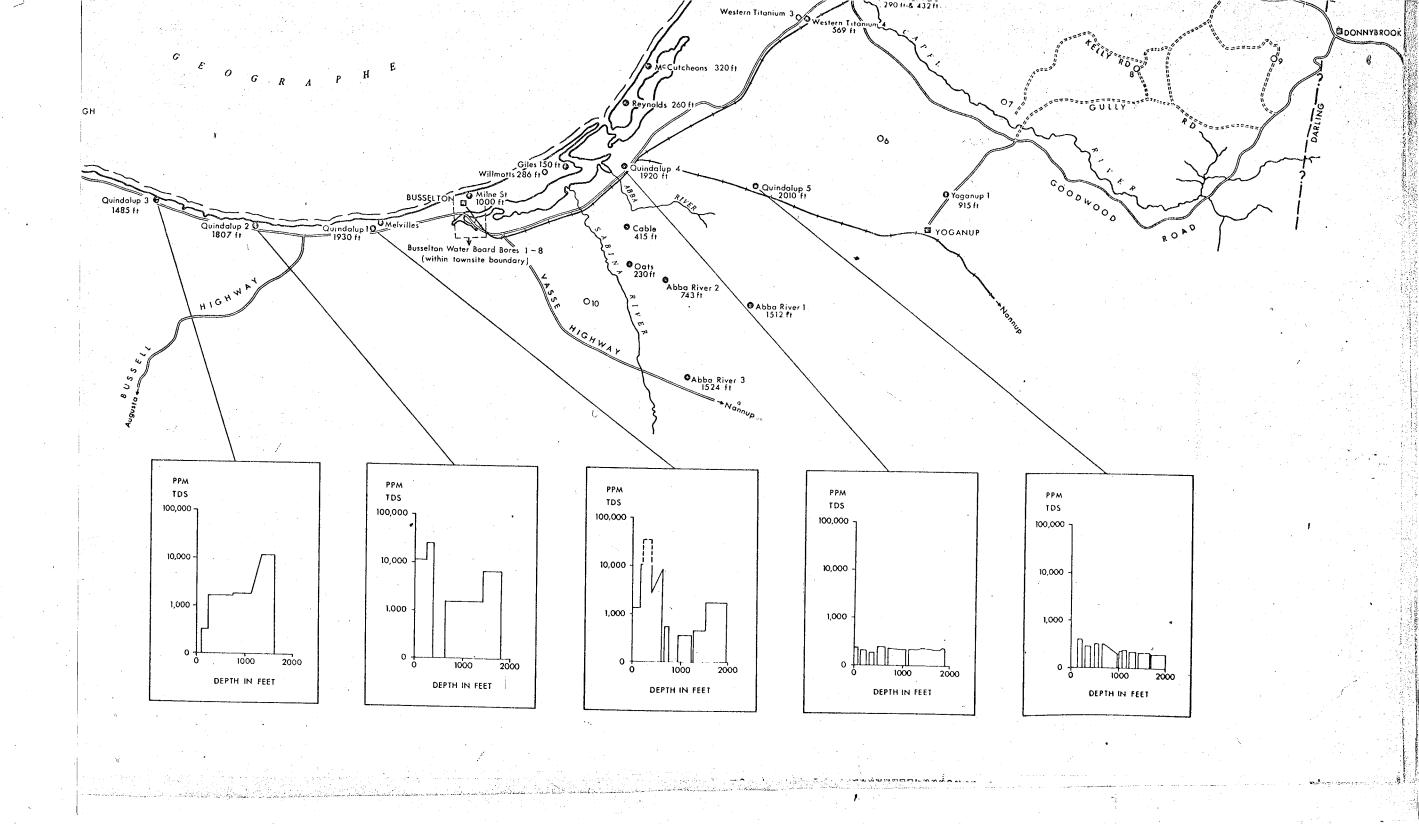
| Bore | | | Depth (feet) | Remarks |
|---|----------|--------|-----------------|-------------------------------|
| G.S.W.A. Quindala | p 1 | | 1930 | 8 in, casing to 504 feet |
| | 2 | | 1807 | Plugged, capped and abandoned |
| | 3 | | 1485 | Plugged, capped and |
| | 4 | **** | 1910 | 10 in. standpipe to 80 ft. |
| | 5 | | 2010 | Plugged, capped and |
| G.S.W.A. Abba Ri | | **** | 1712 | Abandoned |
| | 2 | | 743 1524 | Abandoned - |
| Russelton Water B | 3 | **** | 470 | Drilled 1908, abandoned |
| Busselion water o | HIMERA I | **** | 9/10 | 1960 |
| | 2 | **** | 480 | Abandoned 1938 |
| | 3 | **** | 560(?) | Abandoned |
| | 4 | **.* | 163 | Bore in use |
| | 5 | | 495 | Bore in use |
| | . 6 | **** | 363 | Bore in uso |
| | 7 | | 555(?) | Bore in use |
| | 8 | *** | 490(?) | Bore in uso |
| Milne Street | **** | **** | 1000 351 | Bore in use |
| Western Titanium | 2 | **** | 572 | Bore in use |
| | 3 | **** | 632 | Rore in use |
| | No bore | a 4 dr | | , as an a second second |
| | 5 | | 569 | Bore in use |
| Westralian Oil Car | el 1 | **** | 193 | Bore in use |
| Westralian Oil Yo | zanup 1 | **** | 915 | Bore abandoned |
| | . 2 | | \$61 | Bore abandoned |
| - Western Mineral Sa | nds Cap | el l | 40-5 | Bore in use |
| | | 0 | 493 | Bore equipped |
| Capel town water | supply | 1 | 290 | |
| 45 4 5 | | 2 | 432 520 | Wannamir Dualing and |
| Oats bore | es es | *** | 450 | Wonnerup-Ruabon area |
| Cable (1956) Ruah Cable (1956) Strat | | 4171 | 450 | |
| McCutcheon's bore | | **** | 320 | Wonnerup area |
| Vasse coal bore 1 | | | 159 | 17 |
| 2 | 2 | 1011 | * 143 | |
| · (| 3 | **** | 269 | Drilled 1892-1900 |
| | | | 476 | |
| | | **** | 656 | |
| | 3 | **** | 330 | 14 |
| Blums bore | **** | **** | 297 | Bore in use |
| Reynolds bore | **** | **** | 200 - 286 | Bore in use Bore in use |
| Wilmotts bore | | | 600 | 23VEG 211 USG |

Appendix 2

PALAEONTOLOGICAL DETERMINATIONS IN THE BUSSELTON—CAPEL AREA

Quindalup Bore 1

| feet | |
|---------|--|
| 22 | Quaternary—Recent |
| 110 | Lower Cretaceous : South Perth Formation |
| 140 | Lower Cretaceous: South Perth Formation (?) |
| 270 | Lower Cretaceous-Upper Jurassic : South Perth Forma- |
| | tion |
| 507 | Barren |
| 620-650 | Lower Jurassic: Cockleshell Gully Formation equivalent (?) |
| 926-934 | Barren |
| 1266 | 1) |
| 1580 | |
| 1680 | Lower Jurassic: Cockleshell Gully Formation equiva- |
| 1690 : | lent (?) |
| 1820 | |
| 1900 | |



Quindalup Bore 2

| feet | , 1 | | | | | |
|--------|------|--------|---------------------|-----------------|----------|-----|
| 110 . | | Lower | Cretaceous (possibl | y Uppermost | Jurassic | ?): |
| 310 | | Sout | h Perth Formation | | | |
| 380 - | | J | | | | |
| 500 . | **** | } | | Omntelan (2) | | |
| 630 | | >Lower | Permian-Upper Art | CEEF-HERRIE (1) | | |
| 820 | | Ì | | | | |
| 960 | **** |) | | | | |
| 1190 | | 1 | | | | |
| 1300 | 1, | | | Almaleta n | | |
| 1450 | | Lower | Permian-Lower Ar | LIUSKIND | | |
| 1730 | | | | | | |
| 1 M(H) | | ì | | | | |

Quindalup Bore 3

| 315 | | Lower Cretaceous (Aptian—Neocomian): South F | 'erth |
|---|----|---|-------|
| 640 840 1140 1165-11 1282 1289 1484 | 70 | Early Upper Permian to Lower Permian (Upper A skian?) | rtin- |

Quindalup Bore 4

| feet 60 | Lower Cretaceous-Uppermost Jurass | ic: | South | Perth |
|---------------------------------------|--|--------|-------|-------|
| 310-316 | Formation Upper Jurassic: Yarragadee Formation | 1 | | |
| 405 1098 1240-1250 1620-1627 | Barren Upper Jurassic : Yarragadee Formation Upper Jurassic : Yarragadee Formation Barren | L L | | |

Quindalup Bore 5

| feet | - 1 | |
|--------|-------|--|
| 88 | | Lower Cretaceous: South Perth Formation |
| 90 | 1-11 | Lower Cretaceous: South Pertil Pointagen |
| 202 | | ٠ |
| 343 | 22.17 | Barren |
| 484 | 11/1 | J |
| 556 | | Lower Cretaceous: South Perth Formation |
| 585 | | January Power Powe |
| 620 | | Aptian—marine (?): South Perth Formation Lower Cretaceous: South Perth Formation |
| 1051 | | |
| 1055 | 1.71 | Barren |
| Sludge | | Lower Cretaceous : South Perth Formation |
| 1190 | | Lower Cretaceous: South retth Formation |
| 1337 | **** | Upper Jurassic: Yarragadee Formation |
| 1340 | *** | Station of the state of the sta |
| West | tern | Titanium Capel Bore C3 |

Western Titanium Capet Bore Co

| feet | t. | 1. | | | |
|------|---------|-----------|------------|---------------------|-----------|
| 195 | | 1) | | | |
| 210 | | TTOMAN | Invacaio : | Yarragadee | Formation |
| 244 | 9775 | > o tuber | SHIMONE. | A MITT IN PROPERTY. | |
| 260 | 1111 | 11 | | | |
| 309 | e 1 P * | 1 | | | |

Western Titanium Capel Mining Pit

| kk compa to | a a prosection | | | |
|-------------|---|---------------------------------------|--------------------------|-----------------------|
| feet 30 | Sample taken from r Lower Cretaceous— marginal marine | nining area 30 fee Upper Neocomian | et below gro to Lower | ound level Aptian— |

Milne Street Bore (Busselton)

| fee | t | 1 | |
|-----|------|---------|-----------------------------------|
| 150 | 1400 | Tamer | Cretaceous-Neocomian to Aptian |
| 160 | 1444 | : STOME | Cicuscona accomme |
| 300 | | | |
| 550 | **** | `₹ | |
| 690 | | Lorenz | Cretaceous, probably samples poor |
| 946 | | TONGL | Ciciaccount browns |
| 970 | | ·) | |

Capel Town Bore 1

| 86.608 | | | | |
|--------|------|---|-------------------------------------|--------|
| 36 | | į | Quaternary, probably Pielstocene | |
| 70 | 1107 | 1 | Timer Creteceous-Additi to recommen | marine |
| 270 | | - | Lower Cretaceous-Neocomian | |

Appendix 3

CHEMICAL ANALYSES—BUSSELTON-CAPEL AREA

| Bore | Depth (feet) | pΗ | ppm Evap. | Cond. | NaCl | Hard- ness CaCO _a | Alka- lin- ity | Ca | Mg | Na | ĸ | Fe | нсо, | CO. | 804 | Cl | NO. | Fl | bhtz 810* |
|--|---|--------------------------|---------------------------------|--------------------------------------|------------------------------------|------------------------------------|--------------------------------|---------------------------|--------------------|----------------------------|----------------------------|------------------------------|------------------------|------|-------------------------|-------------------------------|--|-------------------------|----------------------------|
| Melville Quindalup 1 | 15 460-510 935-1025 | 6·9 8·3 | 1330 640 1450 | 1440 17,000 730 1700 | 764 11,000 274 1110 | 579 76 178 | 403 225 205 | 143 24 31 | 54 4 24 | 248 202 484 | 6 18 | 0·1 | 274 250 | | 88 61 90 | 464 166 675 | <1 1 | ND ND | 5 8 |
| Quindalup 2 | 1180-1250 1455-1580 1680-1770 238-310 834-884 | 7·5 8·1 7·3 | 2720 | 19,700 3010 42,000 + 3680 | 18,100 2160 1710 | 216 | 233 | 39 C | ONT 29 ONT | A M 1 933 A M 1 | 24 | <0.1 | SA! 284 SA! | MPL) | 185 | 1310 | 3 | ND | 6 |
| Quindalup 3 | 1670-1700 1670-1700 275-315 1050-1060 1390-1485 | 7·1 7·9 8·3 | 1270 1280 8850 | 8930 8930 1470 1440 8850 | 7650 7650 941 913 7840 | 171 93 1980 | 90 155 65 | 60 34 721 | 5 2 43 7 | 404 442 2840 53 | 18 10 53 | V0·1 V0·1 V0·1 | 110 189 79 61 | | 168 124 634 14 | 571 554 4780 73 | ND <1 | ND ND ND ND | 9 18 9 16 |
| Quindalup 4 Quindalup 5 | 326 530-560 1370-1400 695-763 1057-1105 | 7·1 7·6 7·3 | 200 330 420 520 310 | 220 340 420 540 340 | 120 120 166 194 146 | 66 83 43 74 | 50 158 181 210 122 | 5 20 20 11 15 | 8 4 9 | 96 109 174 75 | 10 20 12 18 | <0.1 <0.1 <0.1 <0.1 | 192 220 256 | **** | 25 21 59 15 | 73 101 118 85 108 | VV | ND ND ND ND | 12 13 16 16 |
| Busselton Water B.5 | 1339-1392 1026-1955 1029-2000 585 | 6.7 6.5 6.8 | 330 270 310 240 | 370 310 350 | 178 159 16: | 71 | 112 85 105 | 17 12 13 12 | 11 10 8 9 | 79 63 84 49 30 | 16 17 16 14 15 | <0.1 <0.1 8 0.1 | 104 128 54 82 | | 14 23 10 | 92 99 95 55 | V1 V1 V1 V1 0.2 | 0.2 0.25 ND ND | 18 18 19 ND 12 |
| Milne Street Giles (Wonnerup) Willmott (Won- | 350-620 750-1000 150* 286* | 6.6 7.6 6.5 6.3 | 190 272 250 380 | 361 260 400 | 148 270 | 64 | 48 43 | 44 14 13 | 7 | 37 52 94 | 9 | 0.1 | 58 52 | | 12 4 19 | 54 90 164 75 | <1 <1 <1 | 0.3 | 24 24 19 |
| nerup) Reynolds (Won- nerup) | 200* | 6.3 7.3 | 200 220 | 230 | 12 | 1 | 67 | 6 | 12 | 53 | 14 | 0.1 | 1 | | 14 | 75 | <1 | 0.3 | 21 |
| McCutcheon (Won- nerup) Oats (Ruabon) Cabie(1958)(Ruabon) Western Titanium | 320 320* 450* 622* | 5·9 6·2 6·7 | 530 | 630 240 240 | 42 13 | 103 | 48 | 10 5 11 | 19 7 7 | 145 55 48 | 9 12 20 | 0·1 0·1 (13) | 110 110 | | 22 13 16 | | \\\\\\\\\\\\\\\\\\\\\\\\\\\\\\\\\\\\\\ | 0·4 0·4 0·3 | 26 17 14 |
| Capel 3 Western Titanium | 569* | 6.8 | į | 240 | 9 | 2 52 | 90 | 11 | 6 6 | 49 | 21 | 0. | 110 | | 16 | | | 0.3 | 13 |
| Capel 4 Capel Town Bore 1 Westralian Oil | 2801* 103* | 6.1 | | 380 340 | 20 | 83 | | 11 | 13 10 | 83 67 | | 18 | 1 58 | | 17 | 125 | <1 | 0.2 | 14 |
| (Capel) Western Mineral | 495* | 7.2 | 210 | 230 | 10 | 1 65 | 72 | 0 | 8 | 42 | - | | | 1 | - | 1 | | 0.3 | 16 |
| Sands Cable (Stratham) Westralian Oll | 879* 915* | 7.0 4.6 | | 250 400 | 11 88 | | | 10 | | | | | D 93 | | * 6 | | | 0.1 | 37 |

Appendix 4

CHEMICAL CHARACTERISTICS OF GROUNDWATER

| Bore Depth (fert) Depth (fert) | The state of the s | | | | | | | | | | | | | | | | |
|--|--|---|--|--|---|---|--|---|-----|---|--|-------------------------|--|--|--|--|--|
| 2 Quindalup 1 935-1025 4-0 10-1 6 20 Western Titan- 622 1-0 2-8 1-3 | So. | Bore | | magnes- ium ratio | adsorption Na* | ual sodium bicar- | No | Bore | | magnes- ium ratio | $\frac{\text{adsorption}}{\sqrt{(\text{Ca}^{++} + \text{Mg}^{++})}}$ | ual sodium bicar- | | | | | |
| | 10 10 10 10 10 10 10 10 10 10 10 10 10 1 | Quindalup I Quindalup I Quindalup I Quindalup I Quindalup I Quindalup 3 Quindalup 3 Quindalup 3 Quindalup 4 Quindalup 4 Quindalup 4 Quindalup 4 Quindalup 5 Milne Street Milne Street Busselton 5 Quindalup 5 | 935-1025 1180-1250 1080-1770 275-315 1050-1080 1300-1485 326 530-560 1370-1400 605-763 550-620 750-1000 407-405 1057-1105 1025-1055 1025-1050 | 4.0 0.8 0.4 7.5 8.0 0.4 3.3 1.5 0.8 3.8 0.9 0.9 | 10·1 16·3 27·5 13·5 20·0 27 3·5 4·0 5·1 11·0 2·4 0·8 2·6 2·9 3·2 3·3 | 0-8 3-0 0-3 3-8 3-0 0-6 0-4 0-9 0-6 | 20 21 22 23 24 25 26 27 28 20 | Western Titan- lum 3 Western Titan- lum 4 Cable Ruabon Oats Giles Western Mineral Sands, Capel Willmotts Reynolds Cable Stratham Westralian OH Capel | 569 | 1·0 1·1 0·7 0·3 1·2 0·7 0·3 0·9 0·8 | 2.8 2.9 2.6 6.2 2.8 2.4 4.6 2.7 8.6 3.4 | 1·3 1·5 0·2 | | | | | |

PROSPECTS FOR UNDERGROUND WATER SUPPLIES FOR CARNAMAH TOWNSHIP

by P. Whincup

ABSTRACT.

In the Carnamah district, groundwater suitable for the town supply is available only from Quaternary sand overlying chert of the Proterozoic Moora Groundwater in the chert is too brackish to be of use.

Exploratory drilling of an area of Quaternary sand adjacent to the present pumping field is recommended.

INTRODUCTION

Carnamah is a northern wheatbelt town on the Geraldton Highway, 188 miles north of Perth. It has a population of about 370.

Water is supplied to the town from a catchment which yields an average 3 million gallons per annum, and from shallow sands at Winchester, 7 miles south of the town. Overabstraction from these sands has resulted in a marked increase in groundwater salinity and it has become necessary to locate additional supplies. The Moora Group, a sequence composed mainly of chert and orthoguartzite which provides water for several towns about 50 to 75 miles south of Carnamah, crops out near Winchester. Accordingly a hydrogeological inabout 50 to 75 miles south of Carnamah, crops out near Winchester. Accordingly a hydrogeological investigation has been made of these rocks in the Carnamah area at the request of the Public Works Department, the primary object being to delineate their extent and assess the groundwater potential. Geological mapping with an associated groundwater census was done between 4th and 13th April, 1967

A more detailed geological investigation of the Moora Group has been completed recently by G. H. Low and L. N. Wall and some of their results are incorporated.

TOPOGRAPHY

TOPOGRAPHY

The major control on the topography is geological. The Mt. Scratch Siltstone and the Permian Nangetty Formation form the flat, low-lying saline areas on the west and probably underlie the Yarra Yarra salt lakes. These flats are succeeded to the cast by a narrow zone of low, rounded hills and ridges of the Moora Group, succeeded in turn by gently undulating granite hills. In general the natural surface rises from west to east. Intermittent saline creeks drain westerly towards the large Yarra Yarra saline system. The average annual rainfall is about 15 inches, and ranges between 4 inches and 29 inches. inches and 29 inches.

GEOLOGY

Rocks exposed are of Archaean, Proterözoic, and Permian age with overlying Quaternary alluvium.

Archaean

The Archaean is represented by a massive, medium-grained, porphyritic granite often well-jointed and intruded by dolerite dykes. Generally it either crops out or occurs at shallow depth and weathers to a coarse feldspathic sand with the typical kaolinitic weathering profile often being absent.

Proterozoic

The Moora Group type section near Coomberdale is subdivided by Logan and Chase into four formations. Near Carnamah no subdivision is possible as only the Coomberdale Chert appears to be present. This is a sequence of chert, chert breceia, and orthoquartzite, often well-bedded and dipping at about 30 degrees to the west. Occasional interbedded siltstone appears towards the west.

The chert contains colonial organisms resembling stromatolites (Collenia) which reflect the original calcareous nature of the sediments. Several caves have also been preserved, one directly south of points 4 and 5, while there are several sink holes in the yellow sand area which trends eastwards from points 15 and 79. (See Plate 3.)

The contact between the Archaean and the over-The contact between the Archaean and the overlying Moora Group is either faulted or, less commonly, unconformable. The Moora Group is intruded by dolerite dykes and its structure is complicated by the north-south faulting shown on Plate 3. Farther south Low and Wall (pers. comm.) have found evidence of northeast-trending shear zones with associated minor multi-directional faults and shear zones. They are postulating similar faulting near Winchester but this has not been shown on either of the accompanying plates.

The chert is white, yellowish and reddish and weathers to yellow clay and sand. Extensive tracts of yellow sand, such as immediately south and east of Winchester, overlie clay and chert. The sand varies in thickness from a few inches to perhaps 20 feet and may have been derived from the chert by illuviation of the clay fraction.

The Mt. Scratch Siltstone is a purple-reddish, well-bedded siltstone which dips very steeply to the west. It may be in faulted contact with the Moora Group. Only one very thin dolerite dyke was seen to intrude it.